Chemical and Structural Analysis of Rocks Using X-ray Fluorescence and X-ray Diffraction Techniques

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Abstract—Rock examinations for mining and commercial use are a vital process to save money and time. A variety of methods and approaches have been used to analyze rocks and among them, X-ray fluorescence (XRF), and X-ray diffraction (XRD) techniques proved to be an accurate method. This research was conducted to evaluate the utility and reliability of XRF and XRD to analyze the major and trace elements of rocks as well as their crystalline structures. Results showed that XRF and XRD techniques are fast and reliable, nondestructive and non-invasive analytical tools for mineral analysis, particularly for rocks. For this purpose, the analysis of 28 samples of rocks, which collected from three different places of Kurdistan region-Iraq, is reported. These places are Haybat Sultan (HS) region in Koya city as well as TaqTaq (TA) and Garmuk (GT) district nearby Koya city. Throughout this analysis, 34 major and trace elements were detected in the rocks collected from HS region, whereas forty major and trace elements were detected in the rocks collected from TA district and GT district. In addition, it was found that the structures of these rocks were crystalline in nature. This was proved through the value of diffracted angle and it was found that the majority of these rocks are oxides. The external view (the appearance) of all collected rocks was also discussed and it was found that each rock consists of a variety of minerals which will be the reason that each rock has a specific color.

Index Terms—Minerals, Rocks, X-ray diffraction, X-ray fluorescence.

I. INTRODUCTION

There are a variety of materials available in nature such as rock, soil, oil (some places in the world), and water. Exploring these materials is a matter of interest due to their importance impacts and possible significant applications. Among these materials, rocks with their different types have attracted a special interest by researchers, as the evaluation of the source of the rock is considered as an important task for the purpose

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of oil exploration (Mirza et al., 2017; Fatah and Mohialdeen, 2016; Mohialdeen, Hakimi and Al-Beyati, 2015; Tissot and Welte, 1984). In addition, rocks' investigation could be used in the assessment of the source of the rock, such as measuring the level of maturity, the types of organic materials content of the rock as well as to define the paleoenvironmental conditions (Hunt, 1996). Comparing to other regions around the world, Kurdistan region of Iraq is a rich region full of various types of materials such as those mentioned above and especially rocks (Hassan et al., 2015). These regions include, for example, mountains and the regions nearby it. However, limited studies have been conducted in the concern of this matter. It has been demonstrated that to build a clear view of the characteristics of the sediment's source area, the composition of the rocks should be investigated (Roser and Korsch, 1988) and to do so, trace elements are extremely important (Lopez et al., 2005). In fact, there are three types of rocks, which are being formed (and destroyed). Earth's first rocks were igneous rocks. These form from molten rock that has cooled and hardened. Second, sedimentary rocks are formed from the shells and skeletons of microscopic sea creatures. Finally, metamorphic rock forms when rocks are squeezed and heated deep under the earth's crust.

It is important to mention that X-ray fluorescence (XRF) technique has become one of the major instruments in diagnostics (Meenaa et al., 2018; Sarrazin et al., 1998; Vaniman et al., 1998; Vaniman, Bish and Chipera, 1991) due to its ability to analyze the chemistry of the sample (Sanchez et al., 2008; Navak and Singh, 2007). Besides to XRF (Cornaby et al., 2001), X-ray diffraction (XRD) technique (Wadsworth and Baird, 1989) is also considered to be the most conclusive technique to investigate the phase composition (Mohammed, 2012) of the matter, particularly rocks (Marinangeli et al., 2015; Cornaby et al., 2001; Vaniman et al., 1998; Vaniman, Bish and Chipera, 1991). Each phase (metal or metal oxide) has its own unique diffraction spectrum and therefore, it is possible to distinguish between compounds, as this technique is sensitive to crystal structure rather than to composition (Loubser and Verryn, 2008; Sanchez et al., 2008). XRD has its own importance in rocks investigation, as this technique has the ability to identify the structure of rocks in terms of crystallinity or amorphousness. Dealing with the rocks and investigating their chemical and crystalline structures are of considerable problem that should be taken into consideration due to the impact of rocks' applications in various field such as in industry and particularly in cement industry. Moreover, the existence of metal and/or metal oxide in these rocks is of importance matter due to the possible future applications in a variety of fields such as in electricity. The main aim for conducting this research is to further go inside rocks and analyze their structures, both chemically and structurally.

II. EXPERIMENTAL DETAILS

The experimental process starts with collecting the rock samples from several Kurdistan region of Iraq, then grinding the samples and finally analyzing the rocks using XRF and XRD techniques.

A. Samples (Rocks) Collecting

The rocks that used in this research have been collected from three different places of Kurdistan region of Iraq. Ten of these samples codes of HS1-HS10 were collected from Haybat Sultan (HS) in Koya city and the other ten samples were collected from TaqTaq (TA) codes of TA1-TA10 that located close to Koya city and finally eight samples codes of GT-1-GT8 were collected from Garmuk (GT), which is also related to TA region.

B. Samples (Rocks) Grinding

The bulk of these rocks was then grinded using Mini mill II ball machine of type Malvern Panalytical, Gurgaon (Panalytical India) to obtain a fine powder.

C. Rocks Analysis

In this paper, the collected rocks were chemically and structurally analyzed using both XRF and XRD techniques as discussed in details in below sections.

XRF Technique

The elemental concentration and their emitted energies in KeV for all collected rocks were determined using energy dispersive XRF (EDXRF). All measurements were carried out under vacuum using a Rigaku NEX CG with RX9, Mo, Cu, and Al targets. Chemical elements of high and low concentration were detected within our collected rocks. The EDXRF K^{α} and K^{β} lines intensities were measured for all rocks' elements, in which the applied voltage increased in general with the required lines energies. The X-ray measuring time was 200 s for the Al target, whereas was 100 s for the other targets. Furthermore, X-rays from the X-ray tube pass through an optional filter on their way to the rocks sample. The sample placed in the chamber and measured by 20 mm diaphragm in vacuum. The X-rays are then transport from the sample to the detector which is electrically cooled. The signal is then processed by electronics and sent to a computer program, where the chemical compositions and intensity versus energy is measured (Abdullah, Chaqmaqchee and Anwer, 2017; Chaqmaqchee and Baker, 2016). A schematic diagram of the typical EDXRF is shown in Fig. 1.

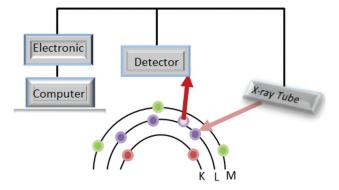


Fig. 1. A schematic of a typical energy dispersive X-ray fluorescence spectrometer.

XRD technique

The crystalline structures of the rocks were analyzed using a high-resolution X-ray powder diffraction technique of type Panalytical Empyrean with CuK^{α} radiation of the wavelength (λ) of 0.15406 nm. The voltage and the current that used throughout the analysis were 45 kV and 40 mA, respectively (Faraj, Kaka and Omar, 2019; Faraj, Chaqmaqchee and Omar, 2017). The XRD (JCPDS card No. 03-1005) data revealed that all the collected rock samples are crystalline in nature.

III. RESULTS AND DISCUSSION

This section has been divided into three sub-sections of XRF, XRD, and the analysis of rocks appearance.

A. XRF

Rigaku NEX CG XRF spectrometer analyses were used to determine the elements of the collected rocks. When these samples are irradiated with X-rays, the intensity as a function of energy can be calculated over the energy range of 1-20 keV at the same computing conditions. Fig. 2 shows the peaks in the ranges of 1.74, 2.32, 3.69, 4.03, 5.89, 6.4, 14.19, 16.83, and 17.48 keV corresponding to Si-k^{α}, S-k^{α}, Ca-k^{α}, Ca-k^{β}, Mn-k^{α}, Fe-k^{α}, Sr-k^{α}, Nb-k^{α}, and Mo- k^{α} lines. It can be seen from Fig. 2a that the intensity of Ca- k^{α} for all HS samples has higher intensity (high Ca concentration) compared to other rocks samples of TA and GT (Figs. 2b and c), whereas the intensity of S- k^{α} , Sr- k^{α} , Mn- k^{α} , and Mo- k^{α} is low or almost negligible. In addition, the intensity of Fe-k^{α} is high (high Fe concentration) in the rocks labeled HS4, HS5, TA1, TA4, TA5, TA9, TA10, and GT7, as shown in Fig. 2. Rocks are divided into two classes according to whether they contain a greater proportion of iron and magnesium or of potassium, sodium, and calcium. Igneous rocks which contain a high percentage of iron and magnesium (ferromagnesian silicates) tend to possess a dark color and are said to be of basaltic or mafic composition. Because of their iron content, such rocks tend to be both denser and darker in color than those rocks composed mainly of non-ferromagnetic silicates. In contrast, igneous rocks

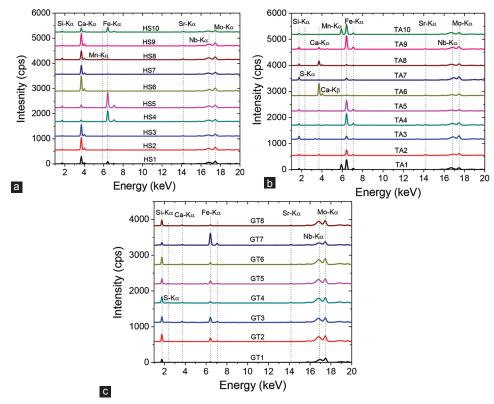


Fig. 2. Intensity versus energy for various rocks of (a) Haybat Sultan, (b) TaqTaq, and (c) Garmuk calculated over the energy range from 1 to 20 keV using X-ray fluorescence with RX9, Mo, Cu, and Al targets.

which contain a greater percentage of potassium, sodium, or calcium (non-ferromagnetic silicates) tend to have a lighter color and are said to be of granitic or felsic composition. Such rocks tend to contain a relatively greater percentage of silica (SiO₂); typically, this is about 70 % by mass. Some minerals are always of the same color, as a consequence of their chemical composition and structure. Their color is a useful aid to identification. Other minerals, such as fluorite, can be different colors due to tiny amounts of chemical impurities, structural flaws, or to their different habits.

Rock types can be classified in terms of chemistry, how the form and environment of formation. Some structures, such as bedding, igneous layering, and gneissic banding, are formed at the same time as the rock itself, but others, such as folding and shearing (stretching of mineral grains), occur later. Many structures exist over a vast range of scale, for example, folding can be seen in microscopic crystal grains as well as across entire cliff faces.

Color is also useful property in rock identification. It can be used to tell apart different varieties of a rock, and it can also provide information on the composition of the rock. In many cases, the color of a rock will be affected by weathering, so a fresh surface should always be exposed.

The ten rock samples collected from HS region have been experienced a geochemical analysis, along with the ten rock samples collected from TA region and the eight rock samples collected from GT region. The raw and processed data are listed in three tables related to the three different places where the rock sample was collected. These tables are labeled as Table I for HS rock samples, Table II for TA rock samples and finally Table III for GT rock samples.

B. XRD

For the purpose of analyzing the crystalline structure of the rocks, XRD technique was used. This technique has been used to analyze the mineralogical composition of the powder materials as well as the phase analysis of the multiphase mixtures. From the values of d-spacing as well as the values of 20, it can be confirmed that the possible minerals of the rock samples could be quartz, alumina, hematite, tridymite, illite and kaolinite, and few other minerals, which is in consistent with Mohammed, 2012; Nayak and Singh, 2007. The main peaks of the XRD data of all rocks are related to Ca-oxide and Fe-oxide such as in HS and TA rocks, whereas they related to silica and Fe-oxide in GT rocks. In addition, alumina (Al₂O₂) is also a major trace in all the three rock samples of HS, TA, and GT. These are in consistent with the XRF data. For example, the main peak of HS1 rock (Fig. 3a) at about 29.6286° can be assigned to diffraction of the (220) plane, which could be for iron oxide. In addition, the XRD pattern of TA rocks exhibits a main plane at about 29.1086°, which are corresponding (hkl) to the (100) plane, as shown in Fig. 4h. Finally, the XRD spectrum of GT rock shows a peak at about 26.4306°, which corresponding to (111) plane of crystalline phase of silica (quartz), as shown in Fig. 5a. The crystallinity of the collected rock samples from the three different places of HS, TA, and GT are proven in the XRD data, as shown in Figs. 3-5. These figures show the similarity

TABLE I
X-ray Fluorescence Measurements for Elements in the Haybat Sultan Rock Samples as Percentages, where the Totals 100%±0.05 Statistical
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						Errors					
No.	Elements				Hayba	t Sultan (HS)	mass concenti	ation (%)			
		HS1	HS2	HS3	HS4	HS5	HS6	HS7	HS8	HS9	HS10
1	CaO	52.7	80.0	77.3	10.6	20.3	91.1	75.3	56.6	75.3	41.5
2	SiO ₂	24.3	9.74	9.80	53.8	41.8	4.63	11.0	37.0	12.1	30.1
3	Al ₂ O ₃	8.15	3.34	3.60	16.8	14.7	1.54	3.50	2.15	3.38	10.8
4	Fe ₂ O ₃	7.25	0.713	0.649	10.4	14.1	0.334	0.906	0.821	0.699	11.0
5	MgO	3.49	5.75	8.11	3.99	4.72	2.01	8.61	1.63	7.74	2.75
6	K,O	2.51	0.194	0.209	2.68	2.59	0.121	0.238	0.149	0.214	1.85
7	TiO,	0.826	0.100	0.118	1.06	1.02	0.0712	0.142	0.0580	0.138	1.10
8	P_2O_5				0.189	0.168	0.0198				0.0968
9	MnO	0.446	0.0266	0.0297	0.193	0.189	0.0214	0.0384	1.44	0.0252	0.329
10	SrO	0.119	0.0323	0.0398	0.0240	0.0401	0.0324	0.0311	0.0551	0.0318	0.0913
11	V ₂ O ₅	0.0522			0.0261	0.0472					0.0242
12	Co ₂ O ₃	0.0229	0.0036	0.0027	0.0331	0.0490	0.0026	0.0054	0.0079	0.0041	0.0521
13	SnO,	0.0181	0.0148	0.0124	0.0152	0.0156	0.0204	0.0131	0.0154	0.0200	0.0209
14	Cr ₂ O ₃	0.0168	0.0055	0.0070	0.0239	0.0285	0.0025	0.0095		0.0173	0.0146
15	NiÔ	0.0158			0.0152	0.0181				0.0081	0.0195
16	ZnO	0.0157	0.0028	0.0031	0.0190	0.0188	0.0025	0.0032	0.0031	0.0098	0.0201
17	CuO	0.0078	0.0042	0.0065	0.0076	0.0134	0.0042	0.0034	0.0038	0.0036	0.0110
18	Rb ₂ O	0.0069	0.0007	0.0007	0.0137	0.0146		0.0011	0.0014	0.0009	0.0164
19	PtO ₂			0.0015							
20	$Y_2O_3^2$	0.0035			0.0042	0.0044			0.0021		0.0057
21	Ga ₂ O ₃	0.0025			0.0027	0.0022					0.0042
22	PbO	0.0018	0.0011		0.0026	0.0032			0.0016		0.0031
23	Ir ₂ O ₃	0.0016			0.0009						0.0014
24	As ₂ O ₃	0.0014	0.0004		0.0044	0.0066			0.0003		0.0036
25	SO ₃		0.0931	0.109	0.0905	0.134	0.0563	0.129	0.0731	0.255	0.0772
26	Ta ₂ O ₅		0.0038	0.0037	0.0031	0.0042		0.0052		0.0033	0.0038
27	BaO				0.0459	0.0398					0.0463
28	Cl				0.0109	0.0085			0.0182	0.0118	
29	HfO ₂				0.0026	0.0029		0.0058			
30	ThO ₂				0.0020	0.0026					
31	TeO				0.0013						
32	Au ₂ Ô				0.0012	0.0012					
33	U_3O_8				0.0007	0.0007					0.0008
34	Nb ₂ O ₅										0.0059

TABLE II

X-ray Fluorescence Measurements for Elements in the TaTaq Rock Samples as Percentages, where the Totals 100%±0.05 Statistical Errors

No.	Elements				Taq	Гаq (TA) mass	concentration	(%)			
		TA1	TA2	TA3	TA4	TA5	TA6	TA7	TA8	TA9	TA10
1	CaO	9.39		4.08	1.96	3.57	90.0	1.35	42.6	12.1	0.197
2	SiO ₂	63.0	24.6	70.5	59.6	29.3	4.79	86.7	49.8	51.9	8.75
3	Al ₂ Õ ₃	16.2	13.4	13.0	21.9	10.4	1.83	7.36	4.46	18.0	1.27
4	Fe ₂ O ₃	6.62	59.6	2.23	8.89	53.2	1.66	1.61	1.50	10.6	88.2
5	MgO	3.10	1.88	2.62	2.75	2.73	0.883	2.09	1.02	3.90	0.638
6	K ₂ O	0.0997		6.94	2.93	0.112	0.188	0.443	0.222	0.267	0.0475
7	TiO ₂	0.670	0.436	0.208	1.23	0.372	0.0800	0.224	0.129	2.30	0.178
8	P_2O_5	0.182			0.143		0.0701			0.322	0.0551
9	MnO	0.184		0.0300	0.0903		0.174	0.0286	0.160	0.268	0.436
10	SrO	0.106		0.0037	0.0315		0.0736	0.0098	0.0280	0.0271	
11	V_2O_5	0.0154	0.0122	0.0078	0.0420	0.0158		0.0073		0.0343	0.0455
12	Co ₂ O ₃	0.0224		0.0097	0.0350		0.0049	0.0094	0.0099	0.0328	
13	SnO ₂	0.0096	0.0050	0.0134	0.0161	0.0025	0.0241	0.0187	0.0126	0.0136	0.0191
14	Cr ₂ O ₃		0.0033	0.0042	0.0191	0.0090	0.0033	0.0136	0.0021		0.0036
15	NiO	0.0031		0.0053	0.0228			0.0105		0.0047	
16	ZnO	0.0097		0.0035	0.0196		0.0061	0.0031	0.0060	0.0157	
17	CuO	0.0043		0.0016	0.0037		0.0061	0.0052	0.0031	0.0071	
18	Rb ₂ O	0.0002		0.0161	0.0194		0.0016	0.0009	0.0012	0.0012	

						TINUED					
No.	Elements	TaqTaq (TA) mass concentration (%)									
		TA1	TA2	TA3	TA4	TA5	TA6	TA7	TA8	TA9	TA10
19	PtO ₂										
20	Y_2O_3	0.0036		0.0039	0.0053		0.0016		0.0032	0.0032	
21	Ga ₂ O ₃	0.0019		0.0018	0.0032			0.0009		0.0015	
22	PbO	0.0006		0.0007	0.0034		0.0021	0.0008	0.0021	0.0011	
23	Ir ₂ O ₃	0.0007			0.0014		0.0019			0.0012	
24	As ₂ O ₃	0.0006		0.0004	0.0012		0.0012	0.0013		0.0008	
25	SO,	0.367	0.0291	0.151	0.0697	0.295	0.214	0.0629	0.0875	0.123	0.0482
26	Ta ₂ O ₅	0.0020		0.0024	0.0022		0.0037		0.0029	0.0040	
27	BaO		0.0080	0.179	0.154			0.0336		0.0205	0.106
28	Cl	0.0066	0.0027	0.0105	0.0109	0.0077	0.0098	0.0152	0.0153	0.0094	0.0077
29	HfO ₂			0.0022						0.0014	
30	ThO ₂	0.0005		0.0017	0.0030					0.0007	
31	TeO2						0.0034	0.0012		0.0019	
32	Au ₂ O	0.0006			0.0007						
33	U_3O_8	0.0006									
34	Nb ₂ O ₅				0.0056					0.0045	
35	WO ₃				0.0013			0.0015			
36	GeO ₂				0.0011						
37	Tl ₂ O ₃				0.0010						
38	HgO				0.0005						
39	Ag ₂ O							0.0006			
40	PdŌ								0.0008		

TABLE II

TABLE III

X-ray Fluorescence Measurements for Elements in the Garmuk Rock Samples as Percentages, where the Totals 100%±0.05 Statistical Errors

No.	Elements			G	armuk (GT) Mass	Concentration (%)		
		GT1	GT2	GT3	GT4	GT5	GT6	GT7	GT8
1	CaO	1.53	1.13	5.58	4.17	2.62	2.36	4.27	5.41
2	SiO ₂	94.5	92.8	85.6	88.8	86.4	93.8	65.1	89.1
3	Al ₂ O ₃	2.40	3.52	4.76	4.72	6.74	2.36	17.9	3.55
4	Fe ₂ O ₃	0.359	1.26	2.14	0.493	1.36	0.658	7.12	0.365
5	MgO	0.762	0.873	1.07	1.24	1.70	0.426	2.35	0.886
6	K ₂ O	0.199	0.199	0.343	0.270	0.570	0.181	1.65	0.197
7	TiO,	0.0524	0.0888	0.236	0.155	0.212	0.0620	0.878	0.0641
8	P_2O_5							0.226	
9	MnO	0.0110	0.0128	0.0549	0.0114	0.249	0.0086	0.155	0.0977
10	SrO		0.0035	0.0091	0.0087	0.0057	0.0054	0.0251	0.0083
11	V ₂ O ₅		0.0044	0.0057		0.0095	0.0033	0.0331	
12	Co ₂ O ₃	0.0020	0.0083	0.0116	0.0029	0.0093	0.0044	0.0330	0.0031
13	SnO ₂	0.0223	0.0091	0.0097	0.0097	0.0103	0.0098	0.0125	0.0115
14	Cr_2O_3	0.0018	0.0064	0.0034	0.0042	0.0038	0.0034	0.0113	0.0016
15	NiO		0.0024	0.0040	0.0019	0.0048	0.0017	0.0068	0.0031
16	ZnO	0.0019	0.0014	0.0080	0.0014	0.0041	0.0011	0.0131	0.0011
17	CuO	0.0018	0.0031	0.0050	0.0015	0.0035	0.0022	0.0020	0.0018
18	Rb ₂ O		0.0005	0.0020	0.0004	0.0022	0.0007	0.0066	0.0005
19	PtO ₂								
20	Y ₂ O ₃		0.0015	0.0017		0.0010	0.0080	0.0044	0.0011
21	Ga ₂ O ₃		0.0004	0.0005		0.0006	0.0002	0.0034	0.0004
22	PbO		0.0006	0.0007	0.0005	0.0007	0.0005	0.0046	0.0025
23	Ir ₂ O ₃								
24	As ₂ O ₃		0.0010			0.0001	0.0003	0.0006	0.0007
25	SO ₃	0.141	0.0502	0.136	0.0480	0.0722	0.0624	0.0532	0.176
26	Ta ₂ O ₅	0.0017		0.0016	0.0011	0.0017	0.0008	0.0033	0.0007
27	BaO			0.0160	0.0268		0.0164	0.0603	0.0662
28	Cl	0.0271	0.0144	0.0181	0.0110	0.0206	0.0055	0.0096	0.0261
29	HfO ₂			0.0012		0.0011	0.0008		
30	ThO ₂							0.0028	

(Contd...)

									1
No.	Elements			Ga	armuk (GT) Mass	s Concentration (%)		
		GT1	GT2	GT3	GT4	GT5	GT6	GT7	GT8
31	TeO ₂	0.0039	0.0013				0.0009		0.0010
32	Au ₂ O								
33	U_3O_8								
34	Nb ₂ O ₅						0.0012	0.0038	
35	WO ₃								
36	GeO ₂	0.0009		0.0005					
37	Tl ₂ O ₃								
38	HgO			0.0004					
39	Ag_2O								
40	PdO								

TABLE III

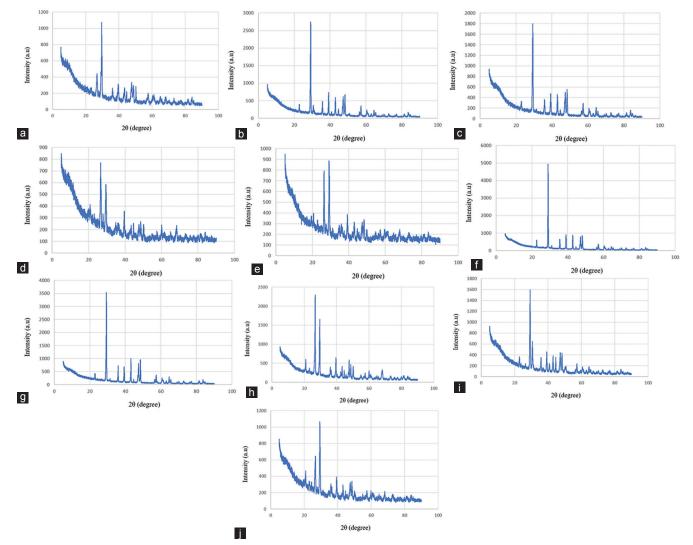


Fig. 3. X-ray diffraction data of the ten different rocks collected from Haybat Sultan place that represented by a-j.

in the XRD data in terms of crystallinity for the majority of the rock samples, although a minor difference can also be seen in these spectra, which could be due the origin place of rocks.

The average crystalline grain size of the powder rock samples was determined using Debye-Scherrer equation (Birks, 1964):

$$D = \frac{0.9}{\beta \cos\theta} \tag{1}$$

where D, λ , β , θ , and 0.9 are the particle size in nm, the wavelength of the X-ray, the full width at half maximum, the peak position, and the Scherrer constant, respectively.

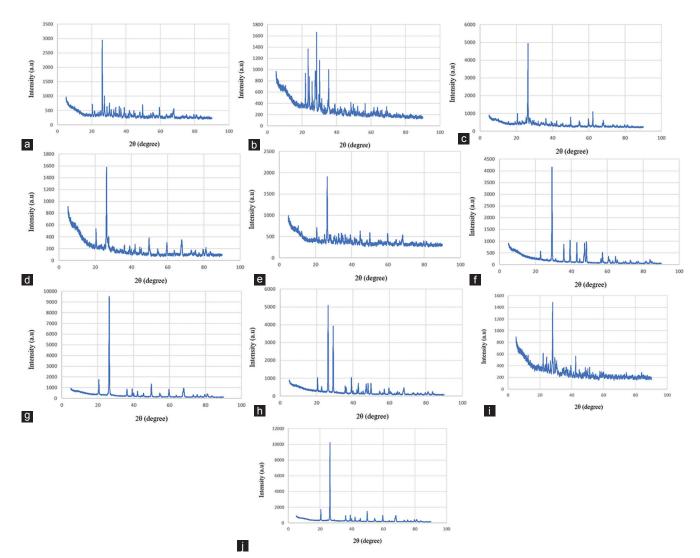


Fig. 4. X-ray diffraction data of the ten different rocks collected from TaqTaq place that represented by a-j.

Using this equation, the crystalline grain size (D) for the collected rocks was found to be as summarized in Table IV:

C. The Rocks Appearance

Rocks on earth can appear with endless array of colors, started from shiny bright and ended with dull dark. The variety of their colors is a result of number of factors. First, the mineral those build up the rocks. Some rock consists only one type of mineral (for example, gold [Au] and silver [Ag]), whereas the majority are consisting of different types of mineral bound together such as ferric oxide (Fe₂O₂) and CaCo₂. Each type of mineral has its chemical composition that determined their unique color. Next factor is the impurities that interact with the minerals and the mineral concentration finally, the atomic bounds within the mineral structure. This factor has observed effect on the absorbed and reflected wavelengths to our eyes, for instance, diamond and graphite are both pure carbon but has different color properties (Mishra, Chhalodia and Tiwari, 2018; Breeding, Magana and Shigley, 2018; Tiwari et al., 2016;

TABLE IV Average Particle Size for the Collected Rocks

Rocks code	Average particle size (nm)
HS	1-4
ТА	2-5
GT	2-7

Sheng et al., 2011). The appearance of the rock samples shows a variety of colors, as represented in Fig. 6. Overall, samples collected from HS contain high concentration of lime (CaO), which can be responsible of their light color (for example, number 6 of Fig. 6a), whereas the gradient in color refers to different concentration of the other mineral oxides, for instance, iron oxide lends the red color to the sample with consideration of their concentration. To study this effect, HS6 (Fig. 6a) and TA6 (Fig. 6b) can be taken which have the same concentration of CaO whereas the ratio of Fe₂O₃ exceeds by 4 times in TA6 compare to HS6. Besides, the shiny appearance of the rock samples, especially GT samples (Fig. 6c) is due to the high content of silica (SiO₂) relative to the other minerals.

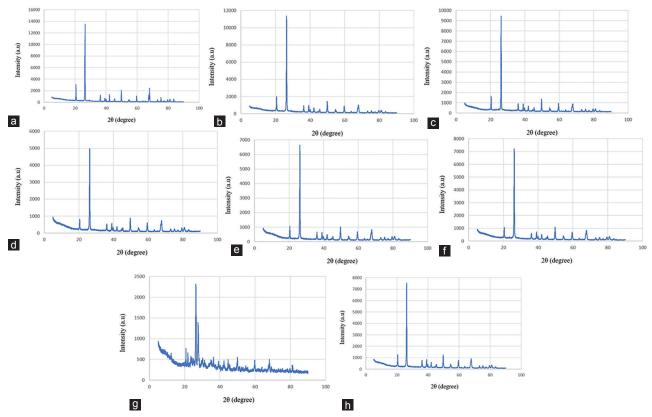


Fig. 5. X-ray diffraction data of the eight different rocks collected from Garmuk place that represented by a-h.



Fig. 6. External view (appearance) of the collected rock samples from (a) Haybat Sultan, (b) TaqTaq, and (c) Garmuk.

IV. CONCLUSION

Twenty-eight samples of different rocks collected from places of Koya city, TA and GT have been analyzed using both XRF and XRD techniques. These two techniques were effective in analyzing the chemical composition and crystalline structure of the collected rocks, as on the one hand, they successfully shown the major and minor minerals in these rocks as well as shown the phases of them, on the other hand. It has been proven from the XRF data that Ca-K^{α}, Fe-K^{α}, and Si- K^{α} were of high concentrations in HS, TA, and GT rock samples, respectively. Furthermore, the XRD data shown that the nature of all collected rocks from HS, TA, and GT were crystalline. The appearance of the collected rocks, on the other hand, was also discussed and it has been shown that the differences in the view of these rocks were due to the various minerals of different concentration contain in these rocks, which gave each rock a specific color.

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